Learning Objectives

Make sure you thoroughly understand the following essential concepts:

- Describe the major features of the *global carbon cycle*, and outline the role of acid-base processes in both the slow and fast parts of the cycle.
- Describe the principal acid-base reactions in *aerobic respiration*.
- Define *acid rain*, and explain its origins. Outline its major effects on the environment.

Acid-base reactions pervade every aspect of industrial-, physiological-, and environmental chemistry. In this unit we touch on a few highlights that anyone who studies or practices chemical science should be aware of.

Some applications of Buffers

Buffer solutions and the buffering effect they produce are extremely important in many practical applications of chemistry. The reason for this is that many chemical processes are quite sensitive to the pH; the extent of the reaction, its rate, and even the nature of the products can be altered if the pH is allowed to change. Such a change will tend to occur, for example, when the reaction in question, or an unrelated parallel reaction, consumes or releases hydrogen- or hydroxide ions.

Buffers in biochemistry and physiology

Many reactions that take place in living organisms fall into this category. Most biochemical processes are catalyzed by enzymes whose activities are highly dependent on the pH; if the local pH deviates too far from the optimum value, the enzyme may become permanently deactivated.

- For example, the oxidation of glucose to carbon dioxide and water — the fundamental energy-producing process of aerobic metabolism — releases H$^+$ ions. The multiple steps of this process are catalyzed by enzymes whose activity is strongly affected by the pH. If the region of the cell where these reactions take place were not buffered, the very act of living would soon reduce the local pH below the range in which the enzymes are active, halting energy production and killing the cell.
- Blood is strongly buffered (mainly by bicarbonate) to maintain its pH at 7.4±0.3; pH values below 7.0 or above 7.8 cause death within minutes.
- The interiors of most cells are buffered near pH 7.0 (by phosphates and proteins)

Buffers in industry

Buffers are employed in a wide variety of laboratory procedures and industrial processes:

- solutions for the calibration of pH meters
- growth of bacteria in culture media
- fermentation processes, including winemaking and the brewing of beer
- controlling the colors of dyes used in coloring fabrics
• shampoos (maintaining pH at or below 7 to counteract the alkalinity of soaps that can cause irritation)
• other personal care products such as baby lotions (keeping the pH around 6 to inhibit the growth of bacteria)
• pH control in printing inks to assure their proper penetration into the paper

Buffers in the environment

• Natural waters (lakes and streams) able to support aquatic life are buffered by the sediments they are in contact with. This buffering is not always able to offset the effects of acid rain.
• Seawater is buffered mainly by bicarbonate and borates. This allows the oceans to absorb about half of the CO$_2$ emitted to the atmosphere by human activities.
• Soils, in order to remain productive must maintain a near-neutral pH, not only for plants but also for microorganisms which fix- and recycle nitrogen. Soils rendered highly acidic by acid rain can keep essential nutrients such as phosphates in forms such as insoluble Ca$_3$(PO$_4$)$_2$ that prevents their uptake by plants.

Acid-base chemistry in physiology

Acid-base chemistry plays a crucial role in physiology, both at the level of the individual cell and of the total organism. The reasons for this are twofold:

• Many of the major chemical components of an organism can themselves act as acids and/or bases. Thus proteins contain both acidic and basic groups, so that their shapes and their functional activities are highly dependent on the pH.
• Virtually all important metabolic processes involve the uptake or release of hydrogen ions. The very act of being alive tends to change the surrounding pH (usually reducing it); this will eventually kill the organism in the absence of buffering mechanisms.

About two-thirds of the weight of an adult human consists of water. About two-thirds of this water is located within cells, while the remaining third consists of extracellular water, mostly interstitial fluid that bathes the cells, and the blood plasma. The latter, amounting to about 5% of body weight (about 5 L in the adult), serves as a supporting fluid for the blood cells and acts as a means of transporting chemicals between cells and the external environment. It is basically a 0.15 M solution of NaCl containing smaller amounts of other electrolytes, the most important of which are HCO$_3^-$ and protein anions.

Respiration, the most important physiological activity of a cell, is an acid-producing process. Carbohydrate substances are broken down into carbon dioxide, and thus carbonic acid:

\[ C(H_2O)_{n} + O_2 \rightarrow CO_2 + n H_2O \]

Maintenance of acid-base balance

It is remarkable that the pH of most cellular fluids can be kept within such a narrow range, given the large number of processes that tend to upset it. This is due to the exquisite balance between a large number of interlinked processes operating at many different levels. Acid-base balance in the body is maintained by two general mechanisms: selective
excretion of acids or bases, and by the buffering action of weak acid-base systems in body fluids.

- Over a 24-hour period, the adult human eliminates the equivalent of about 20-40 moles of H\(^+\) by way of the lungs in the form of CO\(_2\). In addition, the kidneys excrete perhaps 5% of this amount of acid, mostly in the form of H\(_2\)PO\(_4^-\) and NH\(_4^+\). Owing to their electric charges, these two species are closely linked to salt balance with ions such as Na\(^+\) or K\(^+\) and Cl\(^-\).

- The major buffering system in the body is the carbonate system, which exists mainly in the form of HCO\(_3^-\) at normal physiological pH. Secondary buffering action comes from phosphates, from proteins and other weak organic acids, and (within blood cells), the hemoglobin.

There is a huge industry aimed at the notoriously science-challenged "alternative health" market that promotes worthless nostrums such as "ionized water" that are claimed to restore or preserve "the body's acid-base balance". They commonly point out that most foods are "acidic", (i.e., they are metabolized to H\(_2\)CO\(_3\)), but they never explain that most of this acid is almost immediately exhaled in the form of CO\(_2\). The implication is that our health is being ruined by the resulting "acidification" of the body; some further imply this can be a cause of cancer and other assorted ailments.

People who fall for these expensive scams are in effect paying a tax on scientific ignorance. Those who, like you, have studied Chemistry, should consider that they have a social duty to counter this kind of predatory and deceptive pseudoscience.

### Disturbances of acid-base balance

Deviations of the blood plasma pH from its normal value of 7.4 by more than about ±0.1 can be very serious. These conditions are known medically as **acidosis** and **alkalosis**. They can be caused by metabolic disturbances such as diabetes and by kidney failure in which excretion of H\(_2\)PO\(_4^-\), for example, is inhibited.

Numerous other processes lead to temporary unbalances. Thus hyperventilation, which can result from emotional upset, leads to above-normal loss of CO\(_2\), and thus to alkalosis. Similarly, hypoventilation can act as a compensatory mechanism for acidosis. On the other hand, retention of CO\(_2\) caused by bronchopneumonia, for example, can give rise to acidosis. Acidosis can also result from diarrhea (loss of alkaline fluid from the intestine), while loss of gastric contents by vomiting promotes alkalosis.

### Acid rain

The atmosphere is naturally acidic

The natural, unpolluted atmosphere receives acidic, basic, and neutral substances from natural sources (volcanic emissions, salt spray, windblown dust and microbial metabolism) as well as from pollution (Figure 11.6.X). These react in a kind of gigantic acid-base titration to give a solution in which hydrogen ions must predominate to maintain charge balance (indicated by the identical widths of the bar graphs at the bottom labeled "anions" and "cations").
Carbon dioxide, however, is the major source of acidity in the unpolluted atmosphere. As will be explained farther on, CO\(_2\) makes up 0.032 vol-% of dry air, and dissolves in water to form carbonic acid:

\[
\text{CO}_2(g) + \text{H}_2\text{O(l) } \rightleftharpoons \text{H}_2\text{CO}_3(aq)
\]

Thus all rain is “acidic” in the sense that it contains dissolved CO\(_2\) which will reduce its pH to 5.7. The term *acid rain* is therefore taken to mean rain whose pH is controlled by pollutants which can lower the pH into the range of 3-4, resulting in severe damage to the environment.

**Origin of acid rain**

Acid rain originates from emissions of SO\(_2\) and various oxides of nitrogen (known collectively as "NO\(_x\)"") that are formed during combustion processes, especially those associated with the burning of coal. Incineration of wastes, industrial operations such as smelting, and forest fires are other combustion-related sources; these also release particulate material into the atmosphere which plays an important role in concentrating and distributing these acid-forming substances.

**Origins of NO\(_x\)**

Nitrogen oxides are formed naturally by soil bacteria acting on nitrate ions, an essential plant nutrient and itself a product of natural nitrogen fixation. Ammonia, a product of bacterial decomposition of organic matter, is eventually transformed into NO\(_3^-\) by bacteria into nitrogen oxides.

The quantity of fixed nitrogen produced industrially to support intensive agriculture now exceeds the amount formed naturally, and has become the major source of anthropogenic nitrogen oxides.

- Atmospheric nitrogen can be thermally decomposed to its various oxides NO\(_x\) when subjected to high temperatures; one minor natural source is lightning. More significantly, these oxides are also formed when "clean" fuels such as
natural gas (mainly methane, CH₄) is burnt in air.

- NOₓ can also react with other atmospheric pollutants to produce *photochemical smog*.

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### Formation of acids

The gases noted above react with atmospheric oxygen, with each other, and with particulate materials to form sulfuric, nitric, and hydrochloric acids.

- Most of the H₂SO₄ comes from the photooxidation of SO₂ released from the burning of fossil fuels and from industrial operations such as smelting.
- **Nitric acid** similarly results from photooxidation of nitrogen oxides. Large quantities of these oxides are formed in high-temperature combustion processes in gasoline- and jet engines, and in the turbine engines of fossil-fueled electric power plants.
- **Hydrochloric acid** is formed naturally through the reaction of *sea salt aerosols* with atmospheric H₂SO₄ and HNO₃. The salt particles form when ocean waves break the surface and some of the water evaporates before the spray settles back to the surface. On a much smaller scale, fossil fuel combustion and waste incineration are believed to be the principal human-caused sources of HCl.

### Wet and dry deposition

There are two basic mechanisms by which acidic material is transported through the environment. The acid that dissolves in the water droplets that form clouds and precipitation, and is eventually incorporated into fog and snowflakes, undergoes *wet deposition*; this is ordinarily what is referred to as "acid rain". Because it is widely dispersed, often high in the atmosphere, it can travel very large distances from the original source.

About 40% of the acidic substances introduced into the atmosphere becomes adsorbed onto particulate matter, which can be soot, smoke, windblown dust, and salt particles formed naturally from sea spray. This spreads less widely (typically around 30 km from the source), and falls onto surfaces by "dry deposition". Once the acid-laden particles land on surfaces, ordinary rain, fog, and dew release the acidic components, often in considerably more concentrated form than occurs through dry deposition alone.

### Effects on soils

Because soils support the growth of plants and of the soil microorganisms that are essential agents in the recycling of dead plant materials, acid rain has a indirect but profound effect on soil health and plant growth.

↓ Click on image to enlarge
Regions of poorly-buffered soil in the U.S.

Regions of poorly-buffered soil in Canada [gc.ca]

Soils containing alkaline components (most commonly limestone CaCO$_3$ and other insoluble carbonates) can neutralize added acid and mitigate its effects. But soils in high mountain regions tend to be thin and unable to provide adequate buffering capacity. The same is true in almost half of Canada, in which granite rock of the Canadian Shield is very close to the surface; the eastern provinces of the country are strongly impacted by acid rain.

- Most soils also contain clay and humic substances that bind and retain ions such as Ca$^{2+}$, Mg$^{2+}$ and K$^+$ which are essential plant nutrients. Added H$^+$ binds even more strongly to these substances, displacing the nutrient ions from the upper layers. The overall effect is to leach these ions to greater depths where they may be inaccessible to plants, or to wash them away into ground water.
- The SO$_4^{2-}$ component of acid rain converts some nutrient cations into insoluble sulfates, reducing their availability to plants.
- Clays, which are complex aluminosilicates, gradually break down, releasing aluminum ions which normally form insoluble Al(OH)$_3$. Addition of H$^+$ dissolves this hydroxide, raising the concentration of Al(H$_2$O)$_6^{3+}$ to levels that can be toxic to plants.
- Many toxic heavy metals such as Pb, Cd and Cr are present in trace amounts in soils but are tied up in the form of insoluble salts. Acid rain can mobilize the ions of these metals, much as happens with aluminum.

Effects on plants

↓ Click on image to enlarge
The effects on soils noted above affect plants most strongly. However, direct impingment of acidic rain and fog on leaves has other effects which can be especially serious when air pollutants such as SO₂ are present.

- Acid deposition on leaves and needles tends to weaken them by eroding away their protective waxy coatings. This often results in the development of brown spots that interfere with photosynthesis.
- Forests in mountain regions are frequently bathed in clouds and fog which can be even more acidic than normal precipitation, exacerbating the above problem. In addition, the leaves and needles tend to accumulate the fog aerosol into larger droplets. These eventually fall to the forest floor, magnifying the effects on soils described in the previous section.
- Food crops tend to be less affected by acid rain where good farming practices such as the addition of fertilizers to replace depleted nutrients and the addition of limestone to raise the soil pH.

**Effects on lakes and aquatic ecosystems**

Lakes are subject not only to wet and dry acidic deposition, but also by the water they receive from streams and surface runoff. Any toxic elements released by the action of acid rain on soils and sediments can thus be conveyed to, and concentrated within lakes and the streams that empty them.
Regions of poorly-buffered lakes in North America in 1980 [link]

↓ Click on image to enlarge

May Lake in the California Sierras - a typical alpine lake
[link]

Lakes in poorly-buffered areas such as are found in alpine regions (western North America, Colorado and much of Switzerland) or on the Canadian Shield and in the Adirondacs and Appalachians) are highly sensitive to acid deposition.

Severely acidified lakes (such as the one depicted at the above right) can be so devoid of life, including algae, that the water appears to be perfectly clear and bright blue in color.

↓ Click on image to enlarge

pH tolerances of some aquatic organisms
[U.S. EPA]

Aquatic organisms are generally adapted to “ordinary” pH conditions of 6-8, but vary greatly in their tolerance of low pH. As pH declines below 6, the diversity of aquatic animals, plants, and microorganisms diminishes.

- Some plant species such as sphagnum mosses and certain filamentous algae that thrive at very low pH can
proliferate in acidified lakes, producing thick mats that seal off oxygen and thus inhibit the decay of litter on the lake floor.

- Acid deposited onto winter snow is released during spring melting and can cause rapid drops of pH in poorly-buffered lakes that receive waters from these sources. This "spring acid shock" can seriously affect viability of organisms such as fish, amphibians, and insects that lay their eggs in the water which hatch in the spring. The hatchlings are often unable to adapt to the rapid change, and end up deformed or killed.

- Fish damaged by acid deposition [link]

Even those aquatic species able to survive at in low-pH waters can be indirectly affected if their food supply is pH-limited. For example, mayflies and some other insects, which, which are important food sources for frogs, cannot survive below pH 5.5.

- Fish are severely affected by the aluminum that is released from the action of acid on sediments; it causes a coating of mucus to form on the gills, impeding absorption of oxygen. This has led to the extinction of some species from affected lakes.

- Low pH increases the solubility of calcium salts, making it more difficult for bone to form in developing embryos of fish and amphibians.

- Low pH can also make the eggs of aquatic organism thicker and more difficult for embryos to penetrate, thus delaying hatching. When the embryos continue to grow within the confined space of the egg, their spines can be deformed, interfering with their viability when they finally hatch.

Effects on buildings and statues

Acid rain is not new!

"It has often been observed that the stones and bricks of buildings, especially under projecting parts, crumble more readily in large towns where coal is burnt... I was led to attribute this effect to the slow but constant action of acid rain." Robert Angus Smith, 1856

Acid deposition most strongly affects heritage buildings made of limestone and similar carbonate-containing stone materials.

The principal chemical process involves the reaction of sulfuric acid with calcium carbonate:

\[
\text{CO}_3(s) + \text{H}_2\text{SO}_4 \rightarrow \text{CaSO}_4(s) + \text{CO}_2.\]

The acid first erodes and breaks up the surface of the stone. As the hydrated calcium sulfate (gypsum) forms, it picks up iron and other components of the stone and forms an unsightly black coating. Some of this gradually blisters off, exposing yet more stone surface. The gypsum crystals can sometimes grow into the stone, further undermining the surface.
Very old structures such as the Taj Mahal, Notre Dame, the Colosseum and Westminster Abbey have all been affected.

Stone buildings and monuments are rapidly damaged by acid rain [link]

The exposed upper surfaces of the gargoyle are beginning to deteriorate. [link]

Statues and monuments, including those made from marble, are also susceptible to erosion and damage from acid deposition.

Modern buildings are less affected, although acid rain can erode some painted surfaces.

The Geochemical Carbonate system

The carbonate system, consisting as it does of a soluble gas CO₂, soluble ions HCO₃⁻ and CO₃²⁻, and sparingly soluble carbonate salts, spans all four realms of nature: the atmosphere, hydrosphere (mainly the oceans), the lithosphere, and the biosphere. And owing to the acid-base equilibria that govern the transformations between these carbonate species, carbon is readily transported between these geochemical reservoirs. This chapter presents an overview of carbon geochemistry.

Distribution of carbon on Earth

Carbon is the fourth most abundant element in the universe. Within the Earth's crust it ranks 15th, mostly in the form of carbonates in limestones and dolomites. Kerogens, which are fossilized plant-derived carbon mostly in the form of oil shale, constitute another major repository of terrestrial carbon.
The geochemical carbon cycle

The carbonate system encompasses virtually all of the environmental compartments— the atmosphere, hydrosphere, biosphere, and major parts of the lithosphere. The complementary processes of photosynthesis and respiration drive a global cycle in which carbon passes slowly between the atmosphere and the lithosphere, and more rapidly between the atmosphere and the oceans. Thus the "carbon cycle" can be divided into "fast" and "slow" parts, operating roughly on annual and geological time scales.

Another excellent depiction of the carbon cycle, by the U.S. Dept of Energy. (See description):
Carbon Cycle

Carbon dioxide in the atmosphere

$\text{CO}_2$ has probably always been present in the atmosphere in the relatively small absolute amounts now observed. Precambrian limestones, possibly formed by reactions with rock-forming silicates, e.g.

$$\text{CaSiO}_3 + \text{CO}_2 \rightarrow \text{CaCO}_3 + \text{SiO}_2 \label{4-1}$$

have likely had a moderating influence on the $\text{CO}_2$ abundance throughout geological time.

The volume-percent of $\text{CO}_2$ in dry air is 0.032%, leading to a partial pressure of $3 \times 10^{-4}$ $(10^{-3.5})$ atm. In a crowded and poorly-ventilated room, $P_{\text{CO}_2}$ can be as high as $100 \times 10^{-4}$ atm. About $54 \times 10^{14}$ moles per year of $\text{CO}_2$ is taken from the atmosphere by photosynthesis, divided about equally between land and sea. Of this, all except 0.05% is returned by respiration (mostly by microorganisms); the remainder leaks into the slow, sedimentary part of the geochemical cycle where it can remain for thousands to millions of years. Trends in the growth of atmospheric $\text{CO}_2$ for the past five years can be seen on [this US NOAA page](https://www.noaa.gov).

Since the advent of large-scale industrialization around 1860, the amount of $\text{CO}_2$ in the atmosphere has been increasing. Most of this has been due to fossil-fuel combustion; in 1966, about $3.6 \times 10^{15}$ g of C was released to the atmosphere; this is about 12 times greater than the estimated natural removal of carbon into sediments. The large-scale destruction of tropical forests, which has accelerated greatly in recent years, is believed to exacerbate this effect by removing a temporary sink for $\text{CO}_2$.

About 30-50% of the $\text{CO}_2$ released into the atmosphere by combustion remains there; the remainder enters the hydrosphere and biosphere. The oceans have a large absorptive capacity for $\text{CO}_2$ by virtue of its transformation into bicarbonate and carbonate in a slightly alkaline aqueous medium, and they contain about 60 times as much inorganic carbon as is in the atmosphere. However, efficient transfer takes place only into the topmost (100 m) wind-mixed layer, which contains only about one atmosphere equivalent of $\text{CO}_2$; mixing time into the deeper parts of the ocean is of the order of 1000 years. For this reason, only about ten percent of added $\text{CO}_2$ is taken up by the oceans.

Most of the carbon in the oceans is in the form of bicarbonate, as would be expected from the pH which ranges between 7.8 and 8.2. In addition to atmospheric $\text{CO}_2$, there is a carbonate input to the ocean from streams. This is mostly in the form of $\text{HCO}_3^-$ which derives from the weathering of rocks and terrestrial carbonate sediments, and the acid-base reaction

$$\text{H}_2\text{CO}_3 + \text{CO}_3^{2-} \rightarrow 2 \text{HCO}_3^- \label{4-2}$$

which can be considered to be the source of bicarbonate in seawater. In this sense, the ocean can be considered the site of a gigantic acid-base titration in which atmospheric acids (mainly $\text{CO}_2$ but also $\text{SO}_2$, $\text{HCl}$, and other acids of volcanic origin) react with bases that originate from rocks and are introduced through carbonate-bearing streams or in windblown dust.
Carbon in the hydrosphere

Dissolution of CO$_2$ in water

Carbon dioxide is slightly soluble in water:

<table>
<thead>
<tr>
<th>°C</th>
<th>mol L$^{-1}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>0</td>
<td>0.077</td>
</tr>
<tr>
<td>4</td>
<td>0.066</td>
</tr>
<tr>
<td>10</td>
<td>0.054</td>
</tr>
<tr>
<td>20</td>
<td>0.039</td>
</tr>
</tbody>
</table>

Henry's law is followed up to a CO$_2$ pressure of about 5 atm:

\[
[[\text{CO}_2(\text{aq})]] = 0.032 \ P_{\text{CO}_2} \ 
\]

"Dissolved carbon dioxide" consists mostly of the hydrated oxide CO$_2$(aq) together with a small proportion of carbonic acid:

\[
[[\text{CO}_2(\text{aq})]] = 650 \ [[\text{H}_2\text{CO}_3]] \ 
\]

The acid dissociation constant \( K_{a1} \) that is commonly quoted for "H$_2$CO$_3$" is really a composite equilibrium constant that includes the above equilibrium. This means that "pure" H$_2$CO$_3$ (which cannot be isolated) is a considerably stronger acid than is usually appreciated.

Distribution of carbonate species in aqueous solutions

Water exposed to the atmosphere with \( P_{\text{CO}_2} = 10^{3.5} \) atm will take up carbon dioxide, which becomes distributed between the three carbonate species CO$_2$, HCO$_3^-$ and CO$_3^{2-}$ in proportions that depend on \( K_1 \) and \( K_2 \) and on the pH. The "total dissolved carbon" is given by the mass balance

\[
[C_{t}] = [\text{H}_2\text{CO}_3] + [\text{HCO}_3^-] + [\text{CO}_3^{2-}] \ 
\]

The distribution of these species as a function of pH can best be seen by constructing a log \( C \)-pH diagram for \( C_{t} = 10^{-5} \) M.

This Sillén plot is drawn for two different concentrations of the carbonate system. The lower one, in heavier lines, is for a \( 10^{-5} \) M solution, corresponding roughly to atmospheric CO$_2$ in equilibrium with pure water. The upper plot, for a \( 10^{-3} \) M solution, is representative of many natural waters such as lakes and streams where the water is in contact with sediments.

It is important to note that this diagram applies only to a system in which \( C_{t} \) is constant. In a solution that is open to the atmosphere, this will not be the case at high pH values where the concentration of CO$_2$ is appreciable. Under these conditions this ion will react with H$_2$CO$_3$ and the solution will absorb CO$_2$ from the atmosphere, eventually resulting in the formation of a solid carbonate precipitate.
Figure (PageIndex{3}): Sillén plot for carbonate in water.

The lower of the two above plots can be used to predict the pH of $10^5$ M solutions of carbon dioxide, sodium bicarbonate, and sodium carbonate in pure water. The pH values are estimated by using the mass and charge balance conditions for each solution as noted below.

The reasoning leading to these calculations is explained in the discussion of the $10^{-3}$ M carbonate system in a previous chapter.

**Solution of CO\textsubscript{2} or H\textsubscript{2}CO\textsubscript{3}**

$$[H^+] = [OH^-] + [HCO_3^-] + 2 [CO_3^{2-}] \quad (4-6)$$

which, since the solution will be acidic, can be simplified to

$$[H^+] = [HCO_3^-] \quad (point 3) \quad (4-7)$$

**Solution of NaHCO\textsubscript{3}**

$$[H^+] + [H_2CO_3] = [CO_3^{2-}] + [OH^-] \quad (4-8)$$

or

$$[H^+] = [HCO_3^-] \quad (point 3) \quad (4-9)$$

**Solution of Na\textsubscript{2}CO\textsubscript{3}**

$$[H^+] + 2 [H_2CO_3] + [CO_3^{2-}] = [OH^-] \quad (4-10)$$

or

$$[HCO_3^-] = [OH^-] \quad (point 3)$$
Carbon in the oceans

Natural waters acquire carbon from sediments they are in contact with, and of course also from the atmosphere. The pH is an important factor here; CO$_2$ and solid carbonates are more soluble at high pH, and pH also controls the distribution of carbon species, as is seen in the plot just above.

*Carbon content and pH of typical natural waters*

<table>
<thead>
<tr>
<th></th>
<th>rain</th>
<th>river/lake water</th>
<th>sea water</th>
</tr>
</thead>
<tbody>
<tr>
<td>ppm of carbon</td>
<td>1.2</td>
<td>35</td>
<td>140</td>
</tr>
<tr>
<td>pH (unpolluted)</td>
<td>5.6</td>
<td>6.5 - 8.5</td>
<td>7.5 - 8.4</td>
</tr>
</tbody>
</table>

At the slightly alkaline pH of most bodies of water (of which the oceans constitute 97% of the earth's surface waters), bicarbonate is the principal dissolved carbon species. The quantity of organic carbon is fairly small.

*Carbon content of the oceans, moles × 10$^{18}$*

<table>
<thead>
<tr>
<th></th>
<th>CO$_2$(aq)</th>
<th>HCO$_3^-$</th>
<th>CO$_3^{2-}$</th>
<th>dead org.</th>
<th>living org.</th>
</tr>
</thead>
<tbody>
<tr>
<td>moles</td>
<td>0.18</td>
<td>2.6</td>
<td>0.33</td>
<td>0.23</td>
<td>0.05</td>
</tr>
</tbody>
</table>

Contributors

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